Aquatic Viruses and Climate Change

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Abstract

The viral component in aquatic systems clearly needs to be incorporated into future ocean and inland water climate models. Viruses have the potential to influence carbon and nutrient cycling in aquatic ecosystems significantly. Changing climate likely has both direct and indirect influence on virus-mediated processes, among them an impact on food webs, biogeochemical cycles and on the overall metabolic performance of whole ecosystems. Here we synthesise current knowledge on potential climate-related consequences for viral assemblages, virus-host interactions and virus functions, and in turn, viral processes contributing to climate change. There is a need to increase the accuracy of predictions of climate change impacts on virus-driven processes, particularly of those linked to biological production and biogeochemical cycles. Comprehension of the relationships between microbial/viral processes and global phenomena is essential to predict the influence on as well as the response of the biosphere to global change.

Introduction

Whereas most human interventions on ecosystems are immediate, the impact on climate change acts over long periods of time, leading to slight but continuous alterations. As an example, the predicted 2.6–4.8°C warming, under the conditions of Representative Concentration Pathways 8.5 (IPCC, 2013), for this century is expected to be responsible for species movements and extinctions, as well as for changes in the composition of communities, thus likely altering ecosystem functioning (Mooney et al., 2009). Since global change will probably impact all ecosystem components, it is expected that even the smallest life-forms such as Bacteria, Archaea, eukaryotic microorganisms and viruses of aquatic systems will play important roles as agents and recipients of global climate change (Danovaro et al., 2011; Hutchins and Fu, 2017).

Beside direct effects by changes in water temperature, others may occur through altered oceanic circulation along with changing habitats, biogeochemical cycles and food webs. Particularly estuaries and continental shelf regions seem to be affected by changed river runoff due to reduced or enhanced precipitation and flooding. This will result in changes in salinity and nutrients, and in the strength and seasonality of circulation patterns (Danovaro et al., 2011; Cozzi et al., 2012). Changing ocean currents, fluctuations in the depth of the surface mixed layer and stratification might influence light and nutrient availability and ultimately CO₂-dynamics and the carbon cycle (Bauer et al., 2013; Finke et al., 2017). Moreover, warming will increase the extent of the oxygen minimum zones (OMZ), which are a globally important sink for nitrogen and a source of CO₂. Another global player is the reduction of seawater pH due to enhanced CO₂ levels in the atmosphere (ocean acidification, OA). OMZ changes and OA will also influence virus-host interactions.

For inland waters and freshwater systems a whole array of human-mediated threats makes them the most affected ecosystems. Beside pollution and nutrient loads, overexploitation, flow modification, habitat destruction or degradation, climate change is designated as a significant factor for ongoing alterations of these systems (Dudgeon et al., 2006; Zweimüller et al., 2008). There is also evidence that climate warming will affect pelagic carbon metabolism and sediment delivery in lakes and carbon sequestration in streams and rivers (Battin et al., 2009; Boyero et al., 2011; Kritzberg et al., 2014).

Global climate change-related impacts are most striking in the polar regions, where temperature and other factors are changing at more than twice the global average (Hoegh-Guldberg and Bruno, 2010). Due to the major contribution of microbes to ecosystem processes, much attention has been paid to the impact of rapid climate change in Arctic and Antarctic microbial communities (e.g., the European Project on Ocean Acidification, EPOCA). Declines or shifts in these microbial ecosystems will no doubt have implications for entire food webs and biogeochemical fluxes, and polar microbiota can be viewed both as sentinels and amplifiers of global change (Vincent, 2010).

The global relevance of aquatic viruses has been evident for many years. Viruses are the smallest and most numerous biological entities in aquatic ecosystems, the majority being bacteriophages (Wommack and Colwell, 2000; Weinbauer, 2004; Suttle, 2007). They are about 10 to 15-fold more abundant than their microbial hosts, and equivalent to the carbon in ca. 75 million blue whales (ca. 10% of prokaryotic carbon by weight; Suttle, 2005). The typically high frequency of virus infection in aquatic ecosystems is known to be a major cause of prokaryotic host mortality. Host cell lysis in the marine environment was calculated to be a source of up to 10^9 tons of carbon each day, released from the biological pool (Suttle, 2007; Brussaard et al., 2008). If the main control of prokaryotic abundance is via protozoan grazing, most of the carbon will be channelled to higher trophic levels (microbial loop). In contrast, if viral lysis accounts for most prokaryotic losses, then carbon and nutrients are diverted away from larger organisms (Proctor and Fuhrman, 1990; Wilhelm and Suttle, 1999). The biogeochemical consequence of this “viral shunt” (the process of diverting organic carbon away from the grazing food chain via viral lysis of host cells) are changed rates of carbon accumulation in the photic zone (CO_2 release to the atmosphere vs. vertical transport to the meso-/bathypelagic zone). Viruses are also involved in shaping the composition of bacterial communities, either by reducing abundant host taxa or even by introducing new genetic information into their hosts (e.g., Zhang et al., 2007). They are the largest and most diverse genetic reservoir on Earth and may boost the resilience of ecosystems by sustaining multiple species with similar or identical biochemical pathways (Thingstad, 2000; Brussaard et al., 2008; Jacquet et al., 2010; Brum et al., 2015; Guidi et al., 2016; Gregory et al., 2019).

Elucidating the microbial mediation of carbon-cycle feedbacks to climate change appears to be fundamental for understanding ecosystem responses. This

approach would provide a mechanistic basis for carbon-climate modelling (Zhou et al., 2011). Our substantial lack of knowledge explains why the viral compartment is missing from most climate change models. This review tackles the growing evidence that aquatic viruses interact actively with climate alterations and are key biotic components that can influence the microbial feedback on climate change (compare also Danovaro et al., 2011; Mojica and Brussaard, 2014).

Impact on viral abundance, distribution and dynamics

In most natural aquatic environments virus communities were found to be highly dynamic, with often pronounced changes in abundance and diversity over broad ranges of space and time. Viral abundance is largely linked to host availability. The numbers of planktonic viruses commonly range between 10^4 and 10^8 ml⁻¹, being generally higher in inland waters than in marine systems; peak values were reported for very productive estuaries and lakes (Peduzzi and Luef, 2009; Parikka et al., 2017). Exceptionally high virioplankton abundance has been documented in the alkaline-saline lake Nakuru in Kenya, East Africa (up to 7×10^9 ml⁻¹; Peduzzi et al., 2014). Abundance may be up to three orders of magnitude lower in deep oceanic waters (e.g. Parada et al., 2006; Magagnini et al., 2007; Li et al., 2014; Liang et al., 2014; Lara et al. 2017; Muck et al. 2014), but virus-induced relative prokaryotic mortality increases with water depth: below a depth of 1000 m almost all of the prokaryotic heterotrophic production in surface sediment is transformed into organic detritus and dissolved organic matter (DOM). This “viral shunt” releases on a global scale 0.37-0.63 Gt carbon year⁻¹, and is an essential source of labile organic matter for the deep-sea ecosystem (Danovaro et al., 2008). Recently, Lara et al. (2017) estimated that about 145 Gt of C, 27.6 Gt of N, and 4.6 Gt of P are released annually by the viral shunt in global tropical and subtropical oceans. Moreover, on average, 33.6% (equalling 0.605 Pg C year⁻¹) of the globally respired carbon from fluvial systems may pass through a viral loop (Peduzzi, 2016).

Factors affecting virus dynamics and microbial host-virus interactions in the marine environment have been summarized (Mojica and Brussaard, 2014). Typically virus numbers are linked to system productivity, as a result of the balance between production and decay. Eutrophic water bodies contain more virus particles than oligotrophic systems (Parikka et al., 2017). The trophic situation can also be very variable on seasonal scales, particularly in running waters with pronounced hydrological dynamics (Peduzzi, 2016). Virus abundance apparently varies more

pronouncedly on seasonal scales in inland waters than in the marine environment (Wilhelm and Matteson; 2008).

Impact of temperature

The typical viral life cycles (lytic and lysogenic) and replication rates are closely linked with host metabolism. As temperature is a major regulatory factor of microbial growth, an increase in temperature will likely affect the interaction between viruses and their hosts (Mojica and Brussaard, 2014). With increasing prokaryote growth rate, burst size can increase while the length of the lytic cycle decreases (Proctor et al., 1993; Hadas et al., 1997), boosting virus production. Higher temperatures should also increase the contact rates between viruses and hosts (Murray and Jackson, 1992) and hence potentially infection. Although evidence across systems is scarce, increases in burst size together with increases in production have been described for some natural systems (e.g. Parada et al., 2006). Danovaro et al. (2011) suggested that the observed relationship between viral abundance and temperature from different oceanic regions could be used to infer evidence for the potential effect of rising sea-surface temperatures. With these compiled data, the strongest effect was detected in temperate-open oceans: a temperature increase of only a few degrees was accompanied by a doubling in viral abundance. Nonetheless, factors influencing viroplankton distribution are clearly more complex than predicted by temperature alone. This is underlined by the observation that, when global data were grouped together, an overall decreasing pattern of viral abundance with increasing temperature was identified. This inverse relationship suggests that latitudinal changes, which influence radiation regimes and trophic characteristics, can display cascade effects on growth rates of hosts and viral infectivity (Danovaro et al., 2011). A large data set from the Atlantic Ocean revealed that temperature was important in the Sargasso Sea, but not in the northeast Atlantic (Rowe et al., 2008). In experiments over short time scales, changes in temperature did not directly influence viral abundance (Feng et al., 2009).

In the same compilation of literature data, Danovaro et al. (2011) showed that, in cold water systems, higher viral production rates were associated with warmer temperatures, whereas the relationship for systems at tropical and mid-latitudes was less clear. They concluded that different oceanic regions would respond differently to changes in surface temperatures caused by climate change. These temperature-associated changes in viral abundance and production are likely secondary effects

caused by changes in host cell communities and in virus decay rates (Nagasaki and Yamaguchi, 1998; Wells and Deming, 2006). Demory et al. (2017) demonstrated that at suboptimum temperatures, lytic cycle kinetics of a *Micromonas* virus were lengthened and viral yield was reduced. Temperature increase shortened the latent periods, increased the burst size, and affected viral infectivity (Maat et al., 2017). Higher temperature will increase viral decay by enhanced activity of extracellular proteases and nucleases (Suttle and Chen, 1992; Noble and Fuhrman, 1997; Corinaldesi et al., 2010; Wei et al., 2018). In addition, alteration of temperature may affect protein stability and biomolecule elasticity of viral capsid proteins or lipid membranes, and influence the folding and binding of proteins and nucleic acids (Mojica and Brussaard, 2014). An experimental study in two contrasting Arctic systems (Lara et al., 2013) revealed that heterotrophic bacterial and viral abundance, bacterial production and grazing by protists increased at higher experimental temperatures. It remains to be elucidated whether predictable seasonal changes in hosts and viruses are applicable for climate-related long-term and large-scale changes (Danovaro et al., 2011).

The molecular mechanisms and the environmental factors behind the lysogenic decision are still not well understood (Long et al., 2008). The present climate-related changes in environmental conditions (temperature, circulation of water masses, nutrient availability, alterations in system productivity, coastal mixing with freshwater, etc.) may significantly influence viral life strategies. Lysogeny is thought to be favoured under low nutrient concentrations and high virus-to-host cell ratios (Wilson and Mann, 1997). Thus, if nutrient availability becomes limited due to increased vertical stratification caused by rising sea surface temperatures (Sarmiento et al., 2004), a substantial shift to the lysogenic life cycle may be the consequence. Conversely coastal and estuarine environments may shift to increase favouring of the lytic life-style with large-scale implications on heterotrophic and autotrophic organisms (Danovaro et al., 2011). However, Vaqué et al. (2019) found rather the inverse trend related to temperature. Knowles et al. (2016) suggested that lysogeny is also favoured at high bacterial production rates (Piggyback-the-Winner model). As temperature increases will influence growth, there could be also a change in viral life strategies. The direct and indirect effects of temperature on viruses and their interaction with hosts remain to be investigated in more detail.

Effect of changing salinity

Freshening of the oceanic environments due to climate change has been discussed particularly for poleward regions, and freshwater increases of the Arctic Ocean are suggested as being linked to respective changes of the North Atlantic as well (Peterson et al. 2002; Peltier et al., 2006). Changes in salinity can also influence viral lifestyles (Williamson and Paul, 2006), but experimental studies reveal no consistent pattern regarding the lytic-lysogenic-decision (Danovaro et al., 2011 and references therein), whereas Bettarel et al. (2011) found a switch from lytic to lysogenic life styles at high salinity. Recently, viral lysogeny and lysis shifts related to spring-neap tidal cycle were observed in a macrotidal subtropical estuary, which may be caused by freshwater-seawater mixing dynamics (Chen et al., 2019). Either induction or increased phage latent periods have been found under elevated sodium chloride concentrations. There is evidence that salinity can be of importance for the burst size of infected cells. In waters with high salinity, large burst sizes (up to 200) have been described (Guixa-Boixareu et al., 1996). Furthermore, in freshwater environments burst sizes are typically higher than in marine systems (Peduzzi and Luef, 2009). Mixing of these two environments may result in shifts of host growth rates, DOM availability and burst sizes. An interesting experimental study regarding freshwater and seawater mixing (e.g. as in estuaries) demonstrated that production rates of freshwater viruses sharply declined after seawater addition, followed by a rapid (within 48 h) recovery of the viral populations. Conversely, marine viruses were not significantly affected by mixing with freshwater (Cissoko et al., 2008). However, Marine et al. (2013) showed that freshwater prokaryote and virus communities can adapt to a controlled increase in salinity through changes in their structure and interactions. Another study has shown the possibility for freshwater viruses to cross into the marine environment and replicate normally (Sano et al. 2004). These studies and others (Bonilla-Findji et al., 2009; Wei et al. 2019) indicate that viruses can rapidly respond to major shifts in the abundance and community composition of bacterial hosts, which suffer from osmotic shock.

In a meta-analysis Danovaro et al. (2011) showed that systems with lower salinity display higher viral abundance. However, these results could reflect the more specific situation of the sampled areas rather than a general trend because this data set was largely from few estuarine systems. It may also be linked to the different nutrient availability and trophic state of the estuarine environments rather than to a direct

effect of salinity on the viral assemblages. In addition, Parikka et al. (2017) showed in an across-systems study that the ratio of viruses to prokaryote abundance was inversely related to salinity.

Changes in salinity and pH can also influence the extent of virus adsorption to particles (Harvey and Ryan, 2004; Mojica and Brussaard, 2014), thus affecting virus distribution and proliferation. Further, enhanced river runoff to the sea will likely increase the particle concentration in coastal regions. In a riverine environment, for example, the quality, size and age of particles and aggregates, and the exposure time of viruses to aggregates, apparently are the key factors regulating viral abundance (Kernegger et al., 2009; Peduzzi, 2016). Importantly, particles (mainly with organic constituents) appear to play a role as viral scavengers or reservoirs rather than viral factories (Weinbauer et al. 2009). Moreover, adsorption of viruses to suspended particles can stimulate growth of the free-living prokaryotic community, e.g. by reducing viral lysis (Peduzzi and Luef, 2008; Weinbauer et al., 2009). In case salinity and particle load play a significant role in community composition, then the predicted sea-level rise, saltwater and freshwater mixing and temperature changes associated with the present global climate change could influence the success of the dominant viral and host taxa.

In those locations where climate change will increase freshwater input and nutrient concentrations, lytic life strategies could become more important through higher burst sizes, increasing infection rates due to the addition of freshwater viruses, and increased growth rates of prokaryotes and phytoplankton (Danovaro et al. 2011). An accelerated prokaryote-virus production cycle will on one hand release more components (e.g. enzymes) that are known as highly active for viral decay (Wommack and Colwell, 2000). On the other hand, increasing amounts of substrate for prokaryotes from the 'viral shunt' will result in larger burst sizes and potentially enhanced rates of adaptation and evolution. Important information could be drawn from additional experimental, proteomic and genomic data analyses in systems with gradients of salinity and from studies on the lifestyle of viruses and their host interactions. The speed and mechanisms of change, which are currently poorly understood, will be the main controlling factors in the shifts along the freshwater-marine continuum (Danovaro et al., 2011).

Ocean acidification

The uptake of anthropogenic CO₂ in ocean waters is reported to alter the pH (Riebesell et al., 2009). Nonetheless, predictions of the effect of ocean pH changes on the viral compartment are difficult. While the direct effects on marine viruses remain uncertain, it is likely that the most significant changes will be caused by the impact on their host organisms, for example on calcifying protists (Danovaro et al., 2011). Another example is the observed sharp drop in nitrification under increased CO₂ concentrations (Hutchins et al., 2009), suggesting that viruses from ammonia-oxidising bacteria and archaea might be affected. Furthermore, decreased pH can affect a broad range of physiological processes in microorganisms, e.g. those based on a proton gradient across membranes. Moreover, a significant fraction of near-surface prokaryotes was found to carry the pigment proteorhodopsin (which can act as a light-driven proton pump), being quite sensitive to even small decreases in pH (Fuhrman et al., 2008). In the so far largest CO₂ manipulation experiment (European Project on Ocean Acidification, EPOCA), complex impacts on different groups of plankton and on biogeochemical cycles were found (Riebesell et al., 2013, and special volume of Biogeosciences).

The survival, infectivity and adsorption of some bacteriophages were reported to be sensitive to pH (Harvey et al., 2014) and several studies have demonstrated an influence on the proliferation dynamics of phage-host systems (e.g., Larsen et al., 2008; Carreira et al., 2012; Traving et al., 2014; Chen et al., 2015). Highfield et al. (2017) showed that the diversity of EhV was much lower in the high-*p*CO₂ treatment enclosure that did not show inhibition of *E. huxleyi* growth. In the EPOCA experiment, a significant influence of ocean acidification on the relationship between viral and bacterial abundance was found (Brussaard et al., 2013). Ocean acidification did not affect lytic viral production in mesocosms experiments from the Mediterranean Sea and a Norwegian fjord, but lysogeny was stimulated and either linked to phytoplankton production (Tsiola et al., 2017) or bacterial production (Vaqué et al., 2017). In summary, profound effects related to pH-changes on microbial communities and their viruses have been detected, but the topic remains largely unexplored.

Potential impact of climate change on viruses and carbon cycling

Knowledge on the interactive role of viruses in global CO₂ fluxes is largely in its infancy. Four main CO₂ sequestration processes in the ocean are known, the physical

or solubility pump, the carbonate, the biological pump and the microbial carbon pump (Legendre et al., 2015). Conversion of CO₂ into biomass and the subsequent sinking of organic and inorganic components of the plankton in the ocean interior are known as the biological carbon pump. The microbial carbon pump is the conversion of organic matter into forms with low bioavailability (Jiao et al., 2010; Legendre et al., 2015). The effect of rising atmospheric CO₂ on ocean acidification and carbonation is also receiving current attention. Factors affecting these processes are relevant to understand the functioning of the largest ecosystem on Earth, and to better comprehend the global carbon cycle and its implications on climate (Riebesell et al., 2009). It is still a largely open question how and to what extent the ‘viral shunt’ influences the efficiency of the biological pump, the microbial carbon pump and the sequestration of carbon in the ocean interior. After viral infection, a varying but substantial fraction of the released organic material can be more or less efficiently utilized by uninfected heterotrophic prokaryotes, thus entering again the food web. The net effect of the viral shunt is that it detours bacterial production away from being consumed by protists and ultimately converts organic matter into dissolved inorganic nutrients, including respired CO₂ (Wilhelm and Suttle, 1999; Suttle, 2007). Viral particles also make a contribution to the pools and fluxes of DOC, DON and DOP in the global oceans (Suttle, 2005; Jover et al., 2014; Zhang, et al., 2014).

In global carbon cycling models viruses are still practically ignored, making it difficult to come up with comprehensive conclusions. One scenario is that viral lysis increases the efficiency of the biological pump by enriching the proportion of carbon in the sinking particulate material (Suttle, 2007). At the same time, a viral-mediated control on the biological pump will potentially be impacted by climate change in that climate-induced warming of surface oceans could increase the mortality of the pelagic prokaryotes; this would cause an increased release of labile dissolved organic matter (DOM) that can enhance the metabolism and respiration of uninfected cells. Furthermore, a meta-analysis of data from the literature suggested that higher temperatures are associated with exponentially higher decay rates of virioplankton, based on higher levels of exoenzymatic activities (Danovaro et al., 2011). Danovaro et al. (2011) present a conceptual scheme of the impact on the biological carbon pump and of potential feedback mechanisms. In that concept, viral lysis, together with rising CO₂ concentration, cause changes in phytoplankton composition and production. The result would be a modification of both the photosynthetically

produced particulate organic (POC) and inorganic (PIC, by calcifying photoautotrophs) carbon. It would also alter the particle export, in particular the relative ratio of PIC to POC in this material. Thus, the flux of CO₂ between the ocean surface and the atmosphere would be affected as a possible feedback effect on atmospheric CO₂ concentrations. The metabolism of heterotrophic prokaryotic cells could also be enhanced by rising sea-surface temperatures, thus increasing carbon consumption and respiration rates. Two potential scenarios regarding the viral-mediated control on the biological pump were suggested (Danovaro et al. 2011). On one hand, by altering the pathways of carbon fluxes in the sea - when converting living organic matter into dead particulate organic matter and DOM via cell lysis - the viral shunt could have a negative effect on the biological pump. DOM from cell lysis will be retained to a greater extent in surface waters, much of it converted to DIC through respiration and photodegradation. On the other hand, the viral shunt could favour carbon sequestration by increasing the nutrient availability for primary producers. This could increase the amount of living POC in surface waters, thus enhancing carbon export. In addition, other studies highlighted the role of viral lysis for the release of dense and refractory colloidal aggregates (Mari et al., 2005), or for the formation and size increase of organic aggregates (Peduzzi and Weinbauer, 1993), suggesting that viral mortality of host cells can favour carbon transport to the deep sea (viral shuttle; (Sullivan et al., 2017; Yamada et al., 2018). As viruses could also produce positively buoyant aggregates (Weinbauer, 2004), the role of aggregates produced by viruses can be described as viral elevator (Weinbauer et al., 2009). The influence of viral lysis on the stability and formation of aggregates may either increase the retention time of particles in the euphotic zone or increase carbon export to the ocean interior (Weinbauer et al., 2009). However, it is still unclear, whether viruses short-circuit or prime the biological pump (Brussaard et al., 2008). Changes in the amount of carbon release into the deep will no doubt influence also the functioning of deep-sea ecosystems (Smith et al., 2008). Recently, it has been shown that viruses play a more significant role for carbon fixation than previously thought; this influence is mediated by auxiliary metabolic genes (AMGs) (Puxty et al., 2016, 2018). This suggests a closer link of viral infection to the carbon cycle. In a recent review, it has been argued that metabolic reprogramming of infected cells and viral lysis alter nutrient cycling and carbon export in the ocean; however, the net impacts remain uncertain (Zimmerman et al., 2020).

Regarding inland waters the current perception is that the contribution of river networks to fluvial net heterotrophy and CO₂ outgassing has been underestimated in the past (Battin, 2008; Raymond et al., 2013; Acuña et al., 2020). Almost nothing, however, is known about the significance of the viral shunt in flowing inland waters (Peduzzi and Luef, 2008; Peduzzi, 2016). Most fluvial systems harbour and transport high amounts of organic carbon (dissolved and particulate). Although viral lysis may release only a small fraction to this large organic carbon pool, it could still contribute significantly to the readily utilizable carbon (labile and semi-labile fraction) and its remineralisation in fluvial systems. This should be particularly relevant in these systems, which are characterised by a high proportion of allochthonous (terrestrial) aged and recalcitrant carbon (Peduzzi, 2016). Pollard and Ducklow (2011) reported for the Australian Bremer River that terrestrial DOC was partly returned to the atmosphere as CO₂ through bacterial respiration, assisted by bacteriophage lysis of their hosts. This short-circuits the microbial loop. Since streams and rivers are also subjected to climate change-related temperature shifts (Zweimüller et al., 2008), it is likely that the role of viruses in determining the proportion of organic material in horizontal transport and in processing and remineralisation in fluvial waters will be impacted as well.

Conclusions and outlook

The growing knowledge on virus ecology of marine and inland water environments has increased the incorporation of virus-related processes into aquatic food web models. Nonetheless, in climate change models the role of viruses is difficult to assess due to the rudimentary studies and data. Viruses can be anticipated as the ultimate nanoscale drivers/regulators of life in the effect they can have on organisms. Accordingly, aquatic viral ecology will contribute importantly, as the effects of climate and anthropogenic forcing in aquatic food webs are resolved step by step (Brussaard et al., 2008; Zimmerman et al., 2020). Changes in water temperature no doubt affect microbial growth, respiratory rates and carbon assimilation. This makes it extremely important to understand the effect of global warming on microbial communities and their central role in the carbon cycle. Temperature increases are also likely to impact the interactions between viruses and their host cells. Clearly, an improved understanding of viral responses to present climate change would enhance our chances to predict and adapt to potential consequences of such changes (compare Danovaro et al., 2011). Virus infection of microbial cells has the potential for

cascading effects in food webs (Peduzzi et al., 2014). As ocean temperatures rise, precipitation patterns change and freshening of surface oceans occur, simultaneous and intermingled cascading effects may be the result. Global shifts in viral lifestyles in response to changing climate could have dramatic effects on global biogeochemical cycles via cascading effects.

One substantial drawback is the lack of long-term data sets that could be used to identify the relationships between climatic conditions and microbial processes. In a synthesis of the current knowledge and based on relevant case studies, and on meta-analyses of literature data, Danovaro et al. (2011) presented a catalogue of potential consequences and scenarios; some are outlined below together with additional important points:

1. There is some degree of consensus that the effect of temperature on virus related processes will be significant. Effects might be different at different latitudes and oceanic regions (at high latitudes rising temperatures may promote the viral compartment and depress it at the tropics). Increased vertical stratification due to climate change can lead to large-scale nutrient limitations, altering viral life strategies (lytic vs. lysogenic cycle).
2. Changes in salinity and the freshwater-marine continuum will probably impact the abundance, proliferation and life strategies of viruses. They may also influence the geographical success of dominant host taxa and subsequently the success of their viral counterparts. Freshening at the poles may increase the input and spread of freshwater groups of hosts and viruses into marine waters, facilitating enhanced crossing over of marine and freshwater taxa. Apparently, climate change-related impacts are particularly severe in polar regions and the microbiota there might be sentinels of global change.
3. It is currently unclear whether the viral shunt will result in negative or positive effects on the efficiency of the biological pump, and consequently, on the feedback of marine systems on climate. Viruses have the potential to interact with the climate through their contribution to the DOM-pool and to the biogenic particles.
4. The effect of ocean acidification on marine viruses will be largely via the effects of pH on the host organisms. Since some key metabolic processes of microorganisms are sensitive to changes in pH of the medium, this may

profoundly influence the overall functioning of microbial-mediated processes and virus-host interactions.

5. The role of inland waters, in particular streams and rivers, will be largely re-evaluated through the current perception that the contribution of river networks to fluvial net heterotrophy and CO₂ outgassing has been underestimated in the past. Increased virus activity will probably contribute significantly to the readily utilizable carbon (labile and semi-labile fraction) and its remineralisation. In fluvial waters, climate change-related temperature shifts will influence the role of viruses in determining the proportion of organic material in horizontal transport and in processing and remineralisation.

Figure 7.1 Schematic representation of some potentially important impacts of climate change on the virus-host system and potential effects and feedbacks of this biological compartment on some global processes and pools; for details see text.

6. Enhanced river runoff to the sea and mixing processes will likely increase the particle concentration at least in coastal regions; such particles are important factors regulating viral abundance. Aquatic aggregates apparently play a role as viral scavengers or reservoirs rather than viral factories, and adsorption of viruses to suspended particles may stimulate growth of the free-living prokaryotic community, e.g. by reducing viral lysis.

From all these considerations it can be concluded that aquatic viruses will be significantly influenced by climate change and that viruses, in turn, can influence processes contributing to climate change. In Fig. 7.1 some potentially important impacts, effects and relations of virus-host systems, facing a global climate change, are outlined. Under the present scenario of climate change, however, it remains unclear whether viruses will ultimately destabilise or even stabilise the dynamics of the living components in ecosystems and the related biogeochemical cycles. Due to the dearth of data it is currently unpredictable whether virus-related processes will amplify or smooth the impact of climate change on aquatic systems (Danovaro et al., 2011). These limitations inhibit incorporating virus-related processes into current climate models on the necessary spatial and temporal scales at which climate change scenarios respond and interact with this important compartment of the biosphere. This calls for focused research priorities and long-term data sets to enhance our knowledge about the role of aquatic viruses in the present climate change scenario.

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